Sung Yong Kim

Department of Mechanical Engineering Korea Advanced Institute of Science and Technology (KAIST)

(Kim, S. Y. 2010, Cont. Shelf Res. 30, 1639 -1655)

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- Submesoscale processes
 - O(1) Rossby number
 [Ro = U/(fL) = ζ/f]
 - A horizontal scale smaller than the first baroclinic Rossby deformation radius; O(1-10) km
 - Frequently observed as fronts, eddies, and filaments

Simulations on mesoscale and submesoscale grids (SST)



- Submesoscale processes
 - O(1) Rossby number [Ro = U/(fL) = ζ /f]
 - A horizontal scale smaller than the first baroclinic Rossby deformation radius; O(1-10) km
 - Frequently observed as fronts, eddies, and filaments
 - Contribute to the vertical transport of oceanic tracers, mass, and buoyancy and rectify the mixed-layer structure and upper-ocean stratification



e.g., vertical frontal scale secondary circulation

Oceanic processes in time and spatial scales



(Chelton 2001, Dickey et al, RG 2006)

- An observational sensor using electromagnetic waves
 - 3-30 MHz frequency (HFR)
 - Using Doppler shift of backscattered signals of surface gravity waves to estimate the background currents
 - Upper 1 m depth-averaged currents



- An observational sensor using electromagnetic waves
 - 3-30 MHz frequency
 - Using Doppler shift of backscattered signals of surface (N) gravity waves to estimate the background currents
 - Upper 1 m depth-averaged currents
 - Surface radial velocity maps on a polar coordinate





High-frequency coastal radar-derived surface currents off the U.S. West Coast



- A network of high-frequency radars (HFRs) along the coast over 2500 km of US West Coast provides km resolution and hourly surface current maps which cover about 150 km offshore from shoreline.
- Due to low signal-to-noise ratio of satellite remote sensing near coastal regions, coastal surface current maps provided by a large HFR system can provide a useful resource to investigate the submesoscale eddies in a view of statistics and dynamics.

Oceanic processes in time and spatial scales



(Chelton 2001, Dickey et al, RG 2006)

High-frequency coastal radar-derived surface currents off the U.S. West Coast (cascade maps)



(Kim et al, JGR 2011, Kim and Crawford, GRL 2014)

 Velocity potential and stream function



- Velocity potential and stream function
- **Divergence and** normalized vorticity







Normalized vorticity

117°20

Longitude (W)

117°10'

0.4

-0.1

-0.2

-0.3

117°30'



- Velocity potential and stream function
- Divergence and normalized vorticity
- Stretching and shearing deformation rates, and strain rate

$$\varrho = \frac{\partial v}{\partial x} + \frac{\partial u}{\partial y},$$

$$\varsigma = \frac{\partial u}{\partial x} - \frac{\partial v}{\partial y},$$

$$\kappa = \sqrt{\varrho^2 + \varsigma^2}.$$
c)
d)

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(a) divergence; (b) stretching deformation;(c) vorticity; (d) shearing deformation.

- Surface
 - Hourly and O(1) km resolution surface current maps
 - Their kinematic and dynamic quantities
- Subsurface
 - ADCP current profiles
 - Temperature profiles thermocline movements



Outline for rest of the talk...

- Eddy detection using surface current maps
 - Geometry-based technique
- Interpretation of submeososcale eddies
 - Statistics of diameters and Rossby numbers
 - Verification with circulation
 - Horizontal structure of identified eddies
 - Secondary circulation due to drifting submesoscale eddies
- Summary

Eddy detection on HFR surface currents



- Streamlines (nearly closed polygons) are identified with winding angle method.
- Co-centered streamlines are fitted into an ellipse.
- If the center of ellipses in consecutive time steps is within a drifting range (e.g., 1.5 km) with the same rotation, ellipses are considered as a part of an eddy time series. The length of time series is called as persistency.

(Kim, CSR 2011)

Rossby number and size



- Based on 2-year hourly observations.
- About 700 eddies are identified for each rotation
- O(0.5-1) Rossby number at the center of eddies
- 5 20 km size diameter (L) eddies

(Kim, CSR 2010)

Circulation & WO parameter



(Kim, CSR 2010)

Horizontal structure



• V_{Θ} and ζ/f_{c} have similar shapes to the Taylor eddy.

(Kim, CSR 2010)

Frontal-scale secondary circulation: Expectation



Frontal-scale secondary circulation: Expectation



 Paired eddies (a front) move to 'a', thermoclines fluctuate up- and down-ward.

Frontal-scale secondary circulation: Expectation



- Paired eddies (a front) move to 'a', thermoclines fluctuate up- and down-ward.
- On other hand, due to moving paired eddies (front) to 'b', the thermoclines fluctuate down- and up-ward.

Frontal-scale secondary circulation: Data-derived indicator

Surface

- Stream function, velocity potential
- Divergence and vorticity,
- Shearing and stretching deformation rates, and strain rate

$$\delta = \nabla_H \cdot \mathbf{u} = \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y},$$
$$\zeta = \nabla_H \times \mathbf{u} = \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y},$$
$$\varrho = \frac{\partial v}{\partial x} + \frac{\partial u}{\partial y},$$
$$\varsigma = \frac{\partial u}{\partial x} - \frac{\partial v}{\partial y},$$
$$\kappa = \sqrt{\rho^2 + \varsigma^2}.$$

Subsurface

- ADCP Current profiles [**u** = **u**(z, t)]
- Rotational tendency of whole water column: Vertical rotary coefficients

$$\alpha(t) = \frac{-\sum_{m < 0} S(m,t) + \sum_{m > 0} S(m,t)}{\sum_{m < 0} S(m,t) + \sum_{m > 0} S(m,t)},$$

 Vertical movements of thermoclines [T-string data]

Frontal-scale secondary circulation: Surface & Temp.

Frontal-scale secondary circulation: Subsurface

Two events of submesoscale eddies approaching ADCP/T-string

Demography of sub-mesoscale eddies off the USWC

Summary

- Submesoscale eddies off southern San Diego detected from direct estimate of kinematic and dynamic quantities of HFR observations and ADCP: Rossby number of O(0.5-1) and 5-25 km diameter
- Frontal-scale vertical circulation due to drifting eddies undulates thermoclines.
- Available submesoscale observational resources are very sparse and few, but they may enhance our understanding on the submesoscale process studies.